Density and size segregation in deposits of pyroclastic flow

Noriko K. Mitani,¹ Hans-Georg Matuttis,² and Toshihiko Kadono³

Received 30 March 2004; revised 8 June 2004; accepted 2 July 2004; published 5 August 2004.

[1] The main portion in the deposits of pyroclastic flow shows various characteristic features such as density and size segregation of large fragments in the vertical direction. To understand the mechanism which produces these features, we numerically study granular flows on an inclined rough surface with impurities that are different in density and size from the matrix particles. The simulation shows the occurrence of density and size segregation of large impurities, and the impurities finally concentrate at the top and the bottom. Based on these results, we discuss the mechanism of the segregation. It is found that the behavior of impurities is governed by gravity, buoyancy and the viscosity of the randomly moving matrix. INDEX TERMS: 8404 Volcanology: Ash deposits; 8414 Volcanology: Eruption mechanisms; 8450 Volcanology: Planetary volcanism (5480); 8499 Volcanology: General or miscellaneous. Citation: Mitani, N. K., H.-G. Matuttis, and T. Kadono (2004), Density and size segregation in deposits of pyroclastic flow, Geophys. Res. Lett., 31, L15606, doi:10.1029/2004GL020117.

1. Introduction

[2] Pyroclastic flows, commonly accompanying explosive volcanic eruptions, contain gas and grains of various sizes $(10^{-6}-10^1 \text{ m})$ and densities $(500-3000 \text{ kg m}^-)$ which move down the flanks of volcanoes. The main portion in the deposits of pyroclastic flows has various features [Sparks et al., 1973; Sparks, 1976; Wilson, 1980, 1984, 1985; Palladino and Valentine, 1995; Druitt, 1995, 1998; Calder et al., 2000]. One of the characteristic features is that density and size segregation of large fragments in the vertical direction are often observed ("type 2" indicated in Figure 7b of Wilson [1980]); light pumice and dense rock fragments are enriched towards the top and the bottom, respectively, and the fragment size increases with height (reverse grading) for light pumice while it decreases (normal grading) for dense rock. Both gradings are often observed simultaneously. Another important feature is that, sometimes, pumice and rock fragments are concentrated at the top and the bottom, respectively, regardless of their size ("type 3" indicated in Figure 7c of Wilson [1980]).

[3] Much discussion in volcanology has been devoted to the mechanism which controls the grading. It was presumed that large fragments were dispersed in a fine-grained,

fluidized matrix, and that the principal mechanisms of density and size segregation of rock fragments must be buoyancy and gravitational settling in the fluidized and the mobile matrix, respectively [Sparks et al., 1973; Sparks, 1976; Wilson, 1980, 1984, 1985; Druitt, 1995, 1998]. Fluidization occurs when external forces are sufficiently large to support the weight of the matrix grains. As one possible mechanism for fluidization of the fine-grained matrix, the discharge of gas in upward direction had been proposed [Sparks, 1976; Wilson, 1980, 1984, 1985; Roche et al., 2002; Druitt et al., 2004]. It was also suggested that gas-fluidization is important for the reduction of friction [Sparks, 1976; Wilson, 1985; Roche et al., 2002; Druitt et al., 2004]. As the most probable source of gas with upward velocity, either diffusive release from pumice [Sparks, 1978] or break-up of pumice [Wilson, 1980] was proposed. Another mechanism proposed for fluidization of the matrix is sedimentation in aqueous systems [Druitt, 1995, 1998]; the displacement of water induced by the sedimentation of large and dense particles results in water streaming upward. Furthermore, it was demonstrated that the gas-particle suspension within a certain range of vertical gas velocity shows homogeneous fluidization like an aqueous suspension [Druitt et al., 2004].

[4] On the other hand, there are theoretical [Bagnold, 1954; Savage and Lun, 1988], experimental [Thomas, 2000], and numerical [Campbell and Brennen, 1985a; Campbell, 1989; Cleary and Campbell, 1993; Hirshfield and Rapaport, 1997; Straub, 1997] results that the gravitational or initial kinetic energy can be converted into random motion of grains in granular flows on inclined rough surfaces. These indicate that fine particles become mobile also in the absence of gas or water, as long as the energy input in the particle systems is sufficient.

[5] In this paper, we apply this idea to pyroclastic flows and study the behavior of impurities which are different in density and size from those of the surrounding matrix particles on an inclined rough surface. In Section 2, the details of the numerical method and the results are shown. In Section 3, we discuss the mechanism for density and size segregation. Finally we summarize the results in Section 4.

2. Numerical Calculations and Results

[6] We use a two-dimensional Discrete Element Method (DEM) with soft particle model that enables us to treat multi-body collisions [*Matuttis*, 1998]. The basis of the DEM simulations is the numerical integration of the equations of motion which describe the translation and rotation of the particles in the system [*Cundall and Strack*, 1979]. The number of matrix particles is 3953 and their diameter and the density are denoted by $D_{\rm m}$ and $R_{\rm m}$, respectively. The number of impurities is 6 and their diameters $D_{\rm I}$ and

¹Earthquake Research Institute, University of Tokyo, Tokyo, Japan. ²Department of Mechanical and Control Engineering, University of Electro-Communications, Tokyo, Japan.

³Institute for Research on Earth Evolution, Japan Agency for Marine-Earth Science and Technology, Kanagawa 237-0061, Japan.

Copyright 2004 by the American Geophysical Union. 0094-8276/04/2004GL020117\$05.00



Figure 1. A snap shot at (a) $t' = t/\sqrt{D_m/g} = 0$, (b) t' = 190, and (c) t' = 880. The matrix particles are denoted by closed circles. The densities of the impurities R_l/R_m are 0.3 (open circles) and 1.7 (gray circles), respectively. The cases for $D_l/D_m = 5$, 7, and 9 are shown. The angle of the slope is set to be 10 degrees. Horizontal and vertical axes are normalized by D_m .

densities $R_{\rm I}$ are set to be 5, 7, and 9 times $D_{\rm m}$ and 0.3 and 1.7 times $R_{\rm m}$, respectively. As the shape of the bottom, we choose a sine curve with an amplitude 0.5 times $D_{\rm m}$ and a wavelength 33 times $D_{\rm m}$. After the vertical settling (Y-direction), the bottom surface is tilted (10 degrees). The system has periodic boundaries in X-direction; as a particle passes through the right boundary of the control volume, it re-enters through the left boundary with the same velocity and vertical position [Campbell and Brennen, 1985a, 1985b; Campbell, 1989; Cleary and Campbell, 1993; Hirshfield and Rapaport, 1997; Straub, 1997]. It should be noted that, though real pyroclastic flows contain particles in a large size range, we do not treat large size ratios between impurities and matrix particles because it is computationally infeasible due to computer time and storage requirements.

[7] Figures 1a-1c show the numerical result: (a) t' = $t/\sqrt{D_m/g} = 0$, (b) t' = 190, and (c) t' = 880, for time t and gravity constant g. The light and dense impurities are denoted by open circles and gray ones, respectively. The matrix particles are denoted by closed ones. After tilting, the matrix particles move most wildly near the bottom and the top. The middle region corresponds to the "plug flow region" where the shear velocity is almost zero. The light impurities rise and the dense ones sink. The largest impurities rise and sink fastest and they already reach the top and the bottom at t' = 190 while the smallest impurities move most slowly. If the slope ends at this time step, the grading is kept so that our result at t' = 190 corresponds to the observation that the reverse grading of pumice and the normal grading of rock can occur simultaneously ("type 2" in the work of Wilson [1980]). Furthermore, all light impurities have reached the top and all dense ones have

reached the bottom of the flow bed at t' = 880; this corresponds to the observations that pumice and rock fragments can be concentrated at the top and the bottom of the flow bed regardless of the fragment size ("type 3" in the work of *Wilson* [1980]).

3. Discussion

[8] In Figure 2, the normalized position of the impurities $Y_{\rm I}/\underline{D_{\rm m}}$ averaged over our normalized time duration $t' = t/\sqrt{D_m/g}$ of 0.03 is plotted against normalized time. The cases for $D_{\rm I}/D_{\rm m} = 5$, 7 and 9, and $R_{\rm I}/R_{\rm m} = 0.3$ and 1.7 are shown. The position $Y_{\rm I}/D_{\rm m}$ increases with time for particles with density $R_{\rm I}/R_{\rm m} = 0.3$ and decreases with time for those with density $R_{\rm I}/R_{\rm m} = 1.7$. This density segregation can be attributed to the balance between the gravitational force and the "buoyant" force caused by granular pressure which is the sum of stress components due to collisions and random motion [Campbell and Brennen, 1985b; Shishodia and Wassgren, 2001; Huerta and Ruiz-Suarez, 2004]. Figure 3 shows the granular pressure P normalized by $R_{\rm m}gD_{\rm m}$ at t' = 500. Here P is averaged over a horizontal control volume and over normalized time duration of 300. The granular pressure monotonously decreases with height and its negative gradient generates a buoyant force. Therefore, it is expected that since the light impurities are subjected to a force from the granular pressure which is larger than the gravitational force, they rise, and vice versa, dense impurities sink.

[9] Furthermore, in Figure 2, one sees that the larger the size of the impurities is, the larger is their speed for rising or sinking at t' < 600; this should account for the size grading. It is known that impurities are subjected to a granular viscous force from randomly moving matrix particles when the impurities move relative to matrix particles [*Zik et al.*, 1992], and that the viscous force is a function of the impurity size [*Albert et al.*, 1999; *Trujillo and Herrmann*, 2003]. Therefore, the dependence in the rising- and sinking-speeds on the size of the impurities is expected to be caused by the granular viscous force.



Figure 2. Numerical results of the position of the impurities averaged over normalized time duration of 0.03 against normalized time t'. The cases for $D_{\rm I}/D_{\rm m} = 5$, 7, and 9, and $R_{\rm I}/R_{\rm m} = 0.3$ and 1.7 are shown.



Figure 3. Numerical results for the bulk structure of matrix particles at t' = 500 averaged over a horizontal control volume and over normalized time duration of 300: normalized granular pressure and the kinetic energy of random motion per matrix particle.

[10] From the above consideration, we discuss the mechanism of density and size segregation in three dimensions. The impurities are subjected to a force by the pressure gradient $F_{\rm P}$ and the gravitational force $F_{\rm g}$. In addition, they are subjected to the viscous force $F_{\rm v}$ from the randomly moving matrix, where $F_{\rm v}$ is proportional to the size of the impurity $D_{\rm I}$ and the relative velocity $v_{\rm I}$ between the impurity and the matrix [*Zik et al.*, 1992; *Albert et al.*, 1999; *Trujillo and Herrmann*, 2003]. Thus the behavior of the impurities is governed by $F_{\rm P} \sim -(dP/dY)D_{\rm I}^3, F_{\rm g} \sim -R_{\rm I}D_{\rm I}^3g$, and $F_{\rm v} \sim -D_{\rm I}v_{\rm I}$, where the upward direction is taken to be positive. The force balance in the steady state can be expressed as

$$v_{\rm I} \sim \left(-\frac{dP}{dY} - R_{\rm I}g \right) D_{\rm I}^2 \tag{1}$$

Equation (1) indicates that whether the impurity rises or sinks depends on whether the gravitational force is smaller or larger than the gradient of the granular pressure. Moreover, equation (1) indicates that the larger $D_{\rm I}$ is, the larger is the speed for rising or sinking. Thus the reverse grading of light pumice fragments and the normal grading of dense rock fragments as well as density segregation can be realized in granular flows simultaneously.

[11] Finally we discuss the mobility of the matrix particles, which is indicated by the behavior of impurities in Figures 1 and 2. Figure 3 also shows the kinetic energy of the random motion T per matrix particle (sometimes called granular temperature) normalized by $R_mgD_m^3$ averaged over a horizontal control volume. This indicates that the random motion near the bottom is violent and that even in the plug flow (middle) region, the random motion is sufficient for particles to be mobile in spite of the absence of shear velocity. This implies a flow of kinetic energy of the random motion induced on the free above surface and the rough bottom surface into the plug flow region [*Campbell and Brennen*, 1985a]. Therefore, the roughness of the

bottom plays an important role in the fluidization of matrix particles.

4. Summary

[12] We performed a numerical simulation to explain the simultaneous occurrence of reverse grading of light pumice fragments and normal grading of dense rock fragments in the absence of gas and water. Our results suggest that the degree of size segregation increases with time, i.e., with the distance from the vent. After sufficient time, the impurities reach the top or the bottom of the granular bed regardless of their size, and only the density segregation prevails.

[13] Acknowledgments. We thank T. Koyaguchi, H. Kamata, and M. Ohno for helpful discussion and comments. We also thank two anonymous reviewers for helpful comments on the manuscript. N. K. M. is supported by Research Fellowships of the Japan Society for the Promotion of Science for Young Scientists.

References

- Albert, R., M. A. Pfeifer, A.-L. Barabasi, and P. Schiffer (1999), Slow drag in a granular medium, *Phys. Rev. Lett.*, 82, 205–208.
- Bagnold, R. A. (1954), Experiments on a gravity-free dispersion of large solid spheres in a Newtonian fluid under shear, *Proc. R. Soc. London, Ser.* A, 225, 49–63.
- Calder, E. S., R. S. J. Sparks, and M. C. Gardeweg (2000), Erosion, transport and segregation of pumice and lithic clasts in pyroclastic flows inferred from ignimbrite at Lascar Volcano, Chile, *J. Volcanol. Geotherm. Res.*, *104*, 201–235.
- Campbell, C. S. (1989), Self-lubrication for long runout landslides, J. Geol., 97, 653–665.
- Campbell, C. S., and C. E. Brennen (1985a), Chute flows of granular material: Some computer simulations, J. Appl. Mech., 52, 172–178.
- Campbell, C. S., and C. E. Brennen (1985b), Computer simulation of granular shear flows, J. Fluid. Mech., 151, 167–188.
- Cleary, P. W., and C. S. Campbell (1993), Self-lubrication for long runout landslides: Examination by computer simulation, *J. Geophys. Res.*, *98*, 21,911–21,924.
- Cundall, P. A., and O. D. L. Strack (1979), A discrete numerical model for granular assemblies, *Geotechnique*, 29, 47–65.
- Druitt, T. H. (1995), Settling behaviour of concentrated dispersions and some volcanological applications, J. Volcanol. Geotherm. Res., 65, 27– 39.
- Druitt, T. H. (1998), Pyroclastic density currents, in *The Physics of Explosive Eruptions*, edited by J. S. Gilbert and R. S. J. Sparks, *Spec. Publ. Geol. Soc. London*, 145, 51–182.
- Druitt, T. H., G. Bruni, P. Lettieri, and J. G. Yates (2004), The fluidization behaviour of ignimbrite at high temperature and with mechanical agitation, *Geophys. Res. Lett.*, 31(2), L02604, doi:10.1029/2003GL018593.
- Hirshfield, D., and D. C. Rapaport (1997), Molecular dynamics studies of grain segregation in sheared flow, *Phys. Rev. E*, 56, 2012–2018.
- Huerta, D. A., and J. C. Ruiz-Suarez (2004), Vibration-induced granular segregation: A phenomenon driven by three mechanisms, *Phys. Rev. Lett.*, *92*, doi:10.1103/PhysLevLett.92.114301.
- Matuttis, H.-G. (1998), Simulation of the pressure distribution under a twodimensional heap of polygonal particles, *Granular Matter*, 1, 83–91.
- Palladino, D. M., and G. A. Valentine (1995), Coarse-tail vertical and lateral grading in pyroclastic flow deposits of the Latera Volcanic Complex (Vulsini, central Italy): Origin and implications for flow dynamics, *J. Volcanol. Geotherm. Res.*, 69, 343–364.
- Roche, O., M. Gilbertson, J. C. Phillips, and R. S. J. Sparks (2002), Experiments on deaerating granular flows and implications for pyroclastic flow mobility, *Geophys. Res. Lett.*, 29(16), 1792, doi:10.1029/ 2002GL014819.
- Savage, S. B., and C. K. K. Lun (1988), Particle size segregation in inclined chute flow of dry cohesionless granular solids, *J. Fluid Mech.*, 189, 311– 335.
- Shishodia, N., and C. R. Wassgren (2001), Particle segregation in vibrofluidised beds due to buoyant forces, *Phys. Rev. Lett.*, 87, doi:10.1103/ PhysLevLett.87.084302.
- Sparks, R. S. J. (1976), Grain size variations in ignimbrites and implications for the transport of pyroclastic flows, *Sedimentology*, 23, 147–188.
- Sparks, R. S. J. (1978), Gas release rates from pyroclastic flows: An assessment of the role of fluidisation in their emplacement, *Bull. Volcanol.*, 41, 1–9.

Sparks, R. S. J., S. Self, and G. P. L. Walker (1973), Products of ignimbrite eruptions, *Geology*, 1, 115–118.

Straub, S. (1997), Predictability of long runout landslide motion: Implications from granular flow mechanics, *Geol. Rundsch.*, 86, 415–425.

Thomas, N. (2000), Reverse and intermediate segregation of large beads in dry granular media, *Phys. Rev. E*, 62, 961–974.

- Trujillo, L., and H. J. Herrmann (2003), Hydrodynamic model for particle size segregation in granular media, *Physica A*, 330, 519, doi:10.1016/ S0378-4371(03)00621-6.
- Wilson, C. J. N. (1980), The role of fluidization in the emplacement of pyroclastic flows: An experimental approach, J. Volcanol. Geotherm. Res., 8, 231–249.
- Wilson, C. J. N. (1984), The role of fluidization in the emplacement of pyroclastic flows, 2: An experimental results and their interpretation, *J. Volcanol. Geotherm. Res.*, 20, 55–84.

Wilson, C. J. N. (1985), The Taupo eruption, New Zealand II. The Taupo ignimbrite, *Philos. Trans. R. Soc. London, Ser. A*, 314, 229–310.

Zik, O., J. Stavans, and Y. Rabin (1992), Mobility of a sphere in vibrated granular media, *Europhys. Lett.*, 17, 315–319.

T. Kadono, Institute for Research on Earth Evolution, Japan Agency for Marine-Earth Science and Technology, Kanagawa 237-0061, Japan.

H.-G. Matuttis, Department of Mechanical and Control Engineering, University of Electro-Communications, Tokyo 182-8555, Japan.

N. K. Mitani, Earthquake Research Institute, University of Tokyo, Tokyo 113-0032, Japan. (mitani@eri.u-tokyo.ac.jp)